Potential of CLARREO measurements for improving model ensemble based multi-decadal climate prediction: a statistical perspective

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Climate prediction

- multip ensem

> - consi uncerta measu model

project

© Original Artist • IPCC a Reproduction rights obtainable from **IPCC AR4** MULTI-MODEL AVERAGES AND ASSESSED RANGES FOR SURFACE WARMING ©IPCC 2007: WG1-AR4 6.0 5.0 Year 2000 Constant Global surface warming (°C) 20th century 4.0 3.0 2.0 1.0 0.0 B1 A1T B2 A1B -1.02000 2100 1900

Year

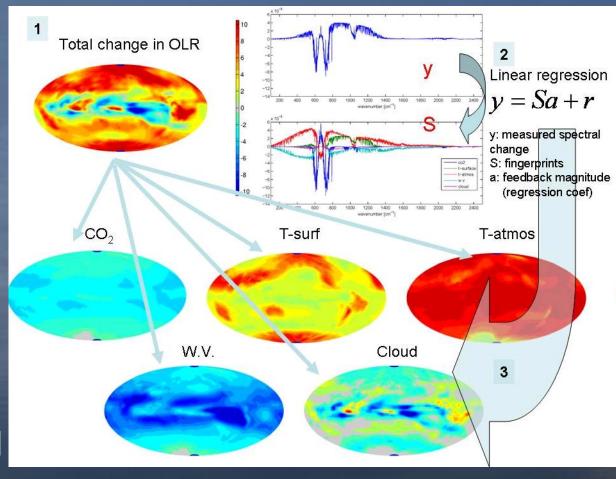
"Tell me, tell me, what are my chances?"

Climate prediction

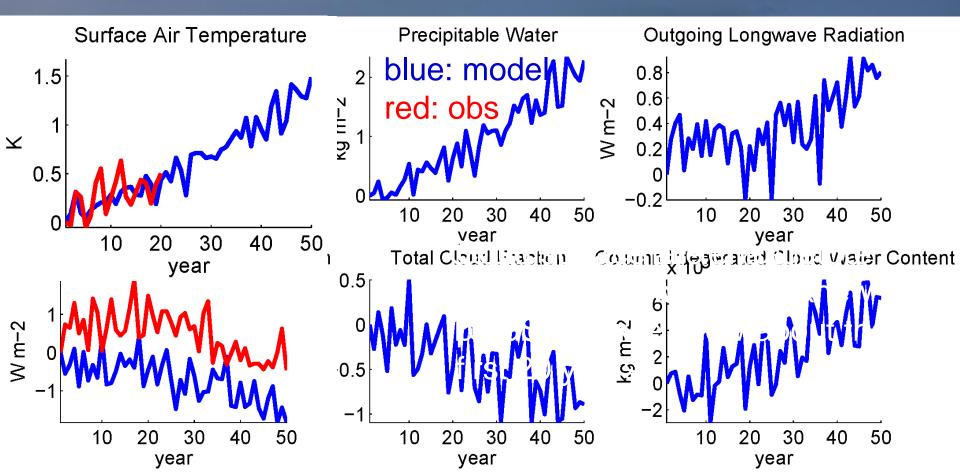
- How might CLARREO help?
 - Through testing and improving the models:
 Observation => model improvement => better prediction based on improved model
 - How to assess the impact of CLARREO NOW?
 - More generally, is it possible to improve the prediction through direct use of data? And what would be optimal data for this purpose?

[Huang, Leroy and Goody, in press, PNAS]

Fingerprinting of the longwave climate feedbacks [Leroy et al., 2008; Huang et al. 2010]



- Model prediction testable hypothesis
- Models predict various variables, a subset of which can be observed.
- Can a climate theory and its prediction of a variable of interest be improved by testing against available observation?



Bayesian Inference

$$P(A | B)P(B) = P(A, B) = P(B | A)P(A)$$



Probability of event A given event B



Probability of event B given event A

Joint probability of events A and B

A: Hypothesis

B: Data

x: hypothetic change that can be measured

d: observation data - change actually measured

$$P(y \mid d) = \int P(x, y \mid d) dx \propto \int P(x, y) P(d \mid x) dx$$

Probability that the prediction of y is true given data

Probability that we would have observed d if the prediction of x was true

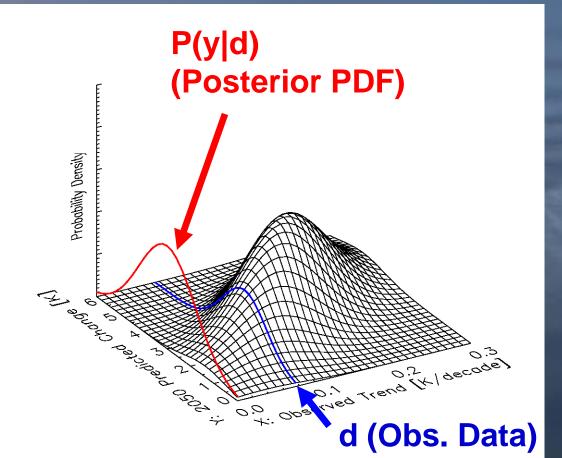
Prerequisite: the relationship between x and y (here, given by the CMIP3 models)

Bayesian Inference

x: hypothetic change that can be measured

d: observation data - change actually measured

$$P(y \mid d) = \int P(x, y \mid d) dx \propto \int P(x, y) P(d \mid x) dx$$



Bayesian Inference

x: hypothetic change that can be measured

Bayesian Method

d: observation data - change actually measured

$$P(y \mid d) = \int P(x, y \mid d) dx \propto \int P(x, y) P(d \mid x) dx$$

Prior estimate

 $y \sim N(\mu_y, \sigma_y^2)$ $x \sim N(\mu_x, \sigma_x^2)$ $d \sim N(d, \sigma_d^2)$

Measurement error

Prior uncertainty in the prediction (model sensitivity difference + natural variability)

Posterior estimate

$$y \mid d \sim N(\mu_{y\mid d}, \sigma_{y\mid d}^2)$$

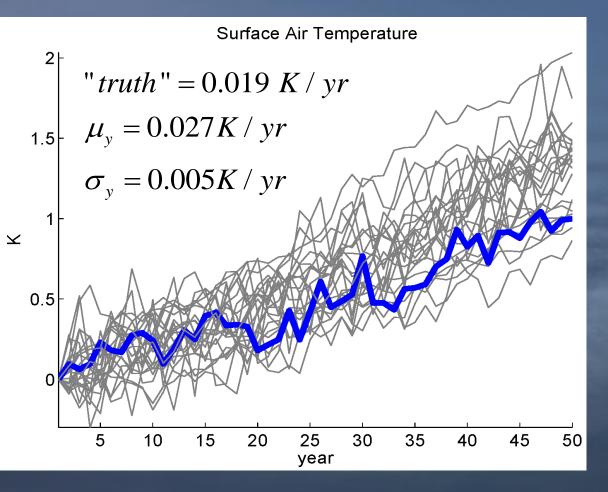
$$\mu_{y|d} = \mu_y + (d - \mu_x) \frac{\sigma_{xy}}{\sigma_x^2 + \sigma_d^2}$$

$$\sigma_{y|d}^2 = \sigma_y^2 (1 - \rho^2 \frac{\sigma_x^2}{\sigma_x^2 + \sigma_d^2})$$

Correlation between x and y



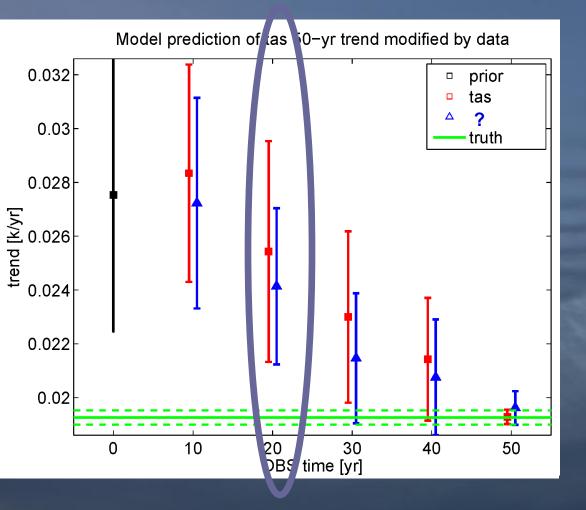
A "perfect model" test



- CMIP3 (IPCC AR4)
 SresA1b experiment
 - One realization each model
 - x, y, d: all linear trends
- One model

 (ncar_pcm1) arbitrarily
 chosen to represent
 observational data,
 the "truth".
- The prediction made by the rest models then validated against this "truth".
- Target prediction: 50year trend in the global mean surface air temperature

A "perfect model" test



$$\mu_{y|d} = \mu_y + (d - \mu_x) \frac{\sigma_{xy}}{\sigma_x^2 + \sigma_d^2}$$

$$\sigma_{y|d} = \sigma_y (1 - \rho^2 \frac{\sigma_x^2}{\sigma_x^2 + \sigma_d^2})^{1/2}$$

- When more and more data are obtained and used to modify the prior prediction according to the above equations, the posterior gets closer and closer to the truth, with less and less uncertainty.
- The best data type would provide the most constraint in the shortest observation time.

Calculations

- Target prediction (y): 50-year trend in surface air temperature
- Observation data (x): 20-year trends in:
 - (In situ) Surface air temperature (Tas), column integrated cloud water and ice, total cloud amount, precipitation, precipitable water (PW), surface downwelling (DLR) and upwelling longwave and shortwave radiation, and atmospheric temperature (Ta), relative humidity, specific humidity (q), and geopotential height (Z) at 500, 200 and 50 hPa levels, and
 - (satellite CLARREO) Outgoing longwave radiation (OLR) and reflected shortwave (RS) radiation at TOA, (clear-sky) spectrally resolved OLR radiances, dry-pressures (P_{dry}) at 5.5, 10 and 14 Km.
- Metrics
 - Improvement in accuracy ($\Delta\mu$) and improvement in precision ($\Delta\sigma$):

$$\Delta \mu = \frac{\mu_{y} - \mu_{y|d}}{\mu_{y} - \mu_{t}} \qquad \Delta \sigma = \frac{\sigma_{y} - \sigma_{y|d}}{\sigma_{y}}$$

In situ

Data Type	ρ	Δσ	Δμ
Z ₅₀₀	0.64	0.22	0.44
Tas	0.61	0.21	0.27
DLR	0.61	0.21	-0.47
PW	0.61	0.18	0.42
Z ₂₀₀	0.56	0.17	0.39
Ta ₅₀₀	0.53	0.15	0.44
q ₅₀₀	0.45	0.10	0.43

Selection criterion: correlation significant at 95% confidence level.

$$\mu_{y|d} = \mu_y + (d - \mu_x) \frac{\sigma_{xy}}{\sigma_x^2 + \sigma_d^2}$$

$$\sigma_{y|d} = \sigma_y (1 - \rho^2 \frac{\sigma_x^2}{\sigma_x^2 + \sigma_d^2})^{1/2}$$

$$\Delta \mu = \frac{\mu_y - \mu_{y|d}}{\mu_y - \mu_t}$$

$$\Delta \sigma = \frac{\sigma_y - \sigma_{y|d}}{\sigma_y - \sigma_{y|d}}$$

Satellite

Data	ρ	Δσ	Δμ	
Type				
OLR	0.67	0.27	0.75	0.8 R ₉₉₅ R ₁₆₀₀
R ₉₉₅	0.67	0.23	0.44	0.6- R ₅₆₀ 1600
P _{dry5.5}	0.64	0.22	0.42	
R ₁₆₀₀	0.60	0.18	0.34	
R ₅₆₀	0.55	0.15	0.35	
RS	-0.47	0.13	0.32	-0.2
R ₅₀₀	0.44	0.09	0.30	-0.4 0 500 1000 1500 2000 2500 3000 cm ⁻¹

Multiple data types used together to improve 50-year Tas trend prediction

Data Type (20-year observation)	Improvement in precision ($\Delta \sigma$)
Tas	0.21
All in situ data	0.35
LW Radiance data only	0.44
All satellite data	0.53

The results here indicate that CLARREO measurements are well chosen for providing powerful constraints on the precision of the ensemble prediction of surface temperature change.

Strict accuracy requirement on observing systems

$$\sigma_{y|d} = \sigma_y (1 - \rho^2 \frac{\sigma_x^2}{\sigma_x^2 + \sigma_d^2})^{1/2}$$

20-yr trends	σ_{x} (inter-model difference)	σ _d (internal variability)	σ _d (WMO GCOS recommendations)			
Tas [K / yr]	0.008	0.002	N/A			
Ta [K / yr]	0.007	0.001	0.005			
PW [kg m ⁻² / yr]		σ ~ 0.033 K	0.07			
DLR [W m ⁻² / yr]	0.044	=> 0.04 mW cm sr ⁻¹ m ⁻² radiance accuracy (at 995 cm ⁻¹ and 280K)				
OLR [W m ⁻² / yr]	0.016	=> ~10 years is re trend accuracy	equired to achieve the			
R ₉₉₅ [mW cm sr ⁻¹ m ⁻² / yr]	0.009	0.003	?CLARREO?			
P _{dry} [hPa / yr]	0.009	0.002	?CLARREO?			

Conclusions/discussions

- We present a methodology that provides constraints on the IPCC multi-model ensemble-based climate prediction by using observation data, and demonstrate that it can be used for selecting optimal data type for this purpose.
- 32 data types are examined for their potential for improving a 50-year surface air temperature trend prediction with data from earlier periods.
- Among the 14 data types that are identified to be of significant potential,
 - The temperature data itself may not be the best data type for constraining surface air temperature prediction;
 - Most constraint is provided by OLR total flux and radiances.
- Given the sample size used to quantify $\rho(x,y)$, confidence on the ranking needs to be substantiated by a large ensemble.

Conclusions/discussions (cont')

- The results indicate that CLARREO measurements are well chosen for constraining ensemble prediction uncertainty and when used together may reduce the uncertainty in the 50-year temperature trend prediction by 50% in 20 years.
- Key to the improvement is the trend measurement accuracy, which constitutes a challenging requirement on most climate observing systems and is where the niche of CLARREO is.
- Yet to answer / improve:
 - Why is, e.g., ρ (OLR,Tas) high?
 - Gaussian assumption for P(x,y)
 - All-sky radiances
 - Optimal combination of data types (radiance selection)
 - Applied to real data



x: hypothetic change that can be measured

d: observation data - change actually measured

$$P(y \mid d) = \int P(x, y \mid d) dx \propto \int P(x, y) P(d \mid x) dx$$

Bayesian Method



Prior estimate

 $y \sim N(\mu_y, \sigma_y^2)$

 $x \sim N(\mu_x, \sigma_x^2)$

 $d \sim N(d, \sigma_d^2)$

Posterior estimate

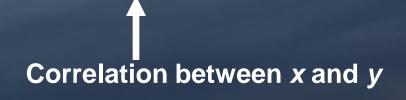


$$\mu_{y|d} = \mu_y + \Sigma_{yx} (\Sigma_{xx} + \Sigma_{dd})^{-1} (d - \mu_x)$$

$$\Sigma_{y|d} = \Sigma_{yy} - \Sigma_{yx} (\Sigma_{xx} + \Sigma_{dd})^{-1} \Sigma_{xy}$$

Measurement error

Prior uncertainty in the prediction (model sensitivity difference + natural variability)



CMIP3 720ppm stabilization (sresA1b) experiment

													[
	model\variable	clivi	clwvi	clt	hur	hus	pr	prw	ps	psl	ta	tas	ts	rlds	rlus	rlut	rsds	rsus	rsdt	rsut	rtmt
NOR		1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	
CAN	cccma_cgcm3_1	5	5	5		5	5	5	5		5	5	5	5	5	5	5	5	5	5	5
CC	cma cgcm3 1 t63	1	1	1		1	1	1			1	1	1	1	1	1	1	1	1	1	1
FRA	_	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	
AUS	csiro_mk3_0	1	1	1	1	1	1	1		1	1	1	1	1	1	1	1	1	1	1	1
	csiro_mk3_5	1	1	1	1	1	1	1	1		1	1	1	1	1	1	1	1	1	1	1
USA	gfdl_cm2_0	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1
	gfdl_cm2_1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1
	giss_aom			2		2	2	2	2	2	2	2	2	2	2	2	2	2	2	2	
	giss_model_e_h	3	3	3	3	3	3	3	3	3	3	3	3			3	3	3	3	3	3
CHN	giss_model_e_r	5	5	5	5	5	3	5	5	5	5	5	5	5	5	5	5	5	5	5	5
ITA	iap_fgoals1_0_g	3	3	3	3	3	3	3	1		3	3	3	3	3	3	3	3	3	3	3
RUS	ingv_echam4			1	1	1	1		1		1	1	1			1				1	1
FRA	C.II.	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1
	TPDI CILI	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1	1
JAP	miroc3_2_hires	1	1	1	1	1	1	1	1		1	1	1	1	1	1	1	1	1	1	1
	miroc3_2_medres	3	3	3	3	3	3	3	3	3	3	3	3	3	3	3	3	3	3	3	3
GER	miub_echo_g		3	3			3	3				3	3	3	3	3	3	3	3	3	3
	mpi_echam5	4	4	4	4	4	4	4	3	3	4	4	3	3	4	4	4	4	4	4	4
JAP	mri_cgcm2_3_2a			1	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5	5
USA	ncar_ccsm3_0	7	7	7	7	7	7	7	6	5	7	7	7	7	6	7	7	7	7	7	7
USA	iicai_pciiii			4	4	4	4	4	1	1	4	4	4	4	4	2	4	4	4	2	2
UK	ukmo_hadcm3	1	1	1	1	1	1		1	1	1	1	1	1	1	1	1	1	1	1	1
OK.	ukmo_hadgem1	1	1	1	1	1	1		1		1	1	1	1	1	1	1	1	1	1	1

surface air temp.

- Source: WCRP CMIP3 multi-model data (ftp-esg.ucllnl.org)
- 24 GCMs
 - -Different number of runs from each GCM
 - -Some diagnostic variables are NOT available from some models
 - –Apparently wrong values in a few cases

Climate sensitivity (50-yr tas trend) of CMIP3 models

mode l	trend	rank	d_trend		int.var.	rank
bccr_bcm2_0	0.02230	5	0.00385	15	0.10747	10
cccma_cgcm3_1	0.02531	8	0.00201	3	0.07267	5
cccma_cgcm3_1_t63	0.02883	16	0.00189	2	0.07255	4
cnrm_cm3	0.03056	21	0.00445	19	0.18647	22
csiro_mk3_0	0.01540	1	0.00294	8	0.11590	12
csiro_mk3_5	0.02980	19	0.00381	14	0.15004	20
gfd]_cm2_0	0.02670	11	0.00256	5	0.12322	16
gfd]_cm2_1	0.02944	18	0.00531	21	0.16838	21
giss_aom	0.02338	6	0.00211	4	0.05419	2 3
giss_model_e_h	0.01976	3	0.00267	6	0.06573	3
giss_model_e_r	0.02029	4	0.00149	1	0.04634	1
iap_fgoals1_0_g	0.02548	. 9 13	0.00392	16	0.19605	23
ingv_echam4	0.02693	12	0.00309	19	0.09537	8
inmcm3_0	0.02982	20	0.00336	12	0.11641	13
ipsl_cm4	0.02809	14	0.00281	7	0.12244	15
miroc3_2_hires	0.04093	24	0.00412	18	0.08123	7
miroc3_2_medres	0.03296 0.02516	23 7	0.00557 0.00645	23 24	0.12773 0.14153	17 19
miub_echo_g	0.02310	13	0.00542	22	0.20434	24
mpi_echam5 mri_cgcm2_3_2a	0.02793	10	0.00342	13	0.20434	14
ncar_ccsm3_0	0.02044	17	0.00334	11	0.07930	6
ncar_ccsiii5_0	0.02929	2	0.00320	10	0.10273	9
ukmo_hadcm3	0.01327	22	0.00319	17	0.12973	18
ukmo_hadgem1	0.02851	15	0.00359	20	0.11078	11
akiio_naageiii1	0.02031	Τ.)	0.00433	20	0.11076	

